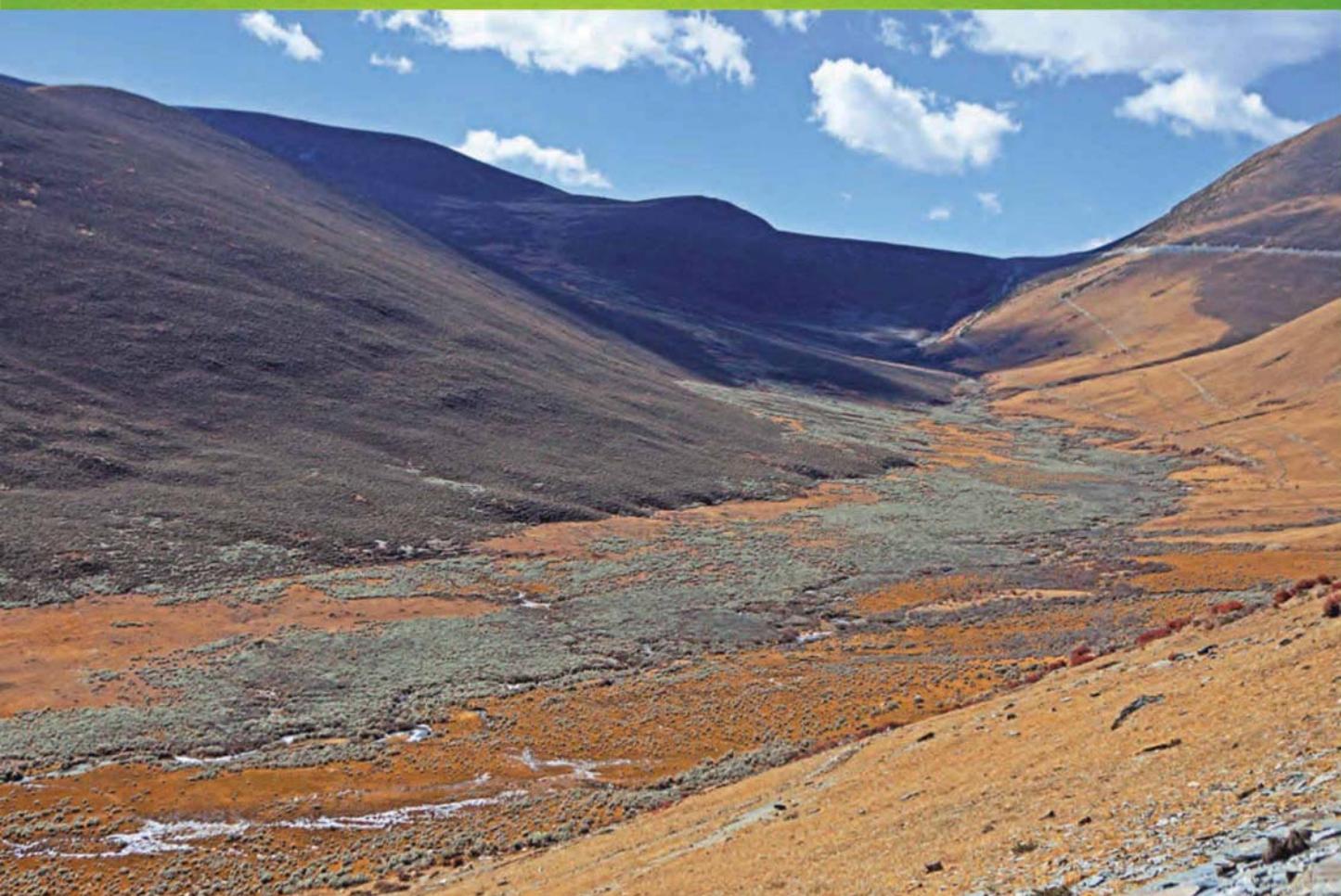


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# 生态学报

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# 生态学报

(SHENTAI XUEBAO)

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**封面图说:** 川西高山地带土壤及植被——青藏高原东缘川西的高山地带坡面上为草地, 沟谷地带由于低平且水分较充足, 生长有很多灌丛。川西地区大约在海拔 4000m 左右为林线, 以下则分布有亚高山森林。亚高山森林是以冷、云杉属为建群种或优势种的暗针叶林为主体的森林植被。作为高海拔低温生态系统, 高山-亚高山地带土壤碳被认为是我国重要的土壤碳库。有研究表明, 易氧化有机碳含量与海拔高度呈显著正相关, 显示高海拔有利于土壤碳的固存。因而, 这里的表层土壤总有机碳含量随着海拔的升高而增加。

彩图及图说提供: 陈建伟教授 北京林业大学 E-mail: cites.chenjw@163.com

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Wang X, Yan Y C, Yan R R, Yang G X, Xin X P. Effect of rainfall on the seasonal variation of soil respiration in Hulunbe Meadow Steppe. Acta Ecologica Sinica, 2013, 33(18): 5631-5635.

## 降雨对草地土壤呼吸季节变异性的影响

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**摘要:**利用土壤碳通量自动观测系统(LI-8150)对呼伦贝尔草原在自然降雨条件下的土壤呼吸作用进行了野外定位连续观测,研究结果表明:降雨对土壤呼吸作用存在激发效应和抑制效应,降雨发生后1—2 h内土壤呼吸速率可增加约1倍,当单次或者连续降雨累积量大于7—8 mm,或土壤含水量大于29%—30%时,降雨对土壤呼吸会产生明显的抑制作用。土壤呼吸的激发效应往往体现在次日,表现为次日平均土壤呼吸速率的显著升高;而抑制效应则在当日即可体现出来,表现为观测当日平均土壤呼吸速率的明显下降。土壤呼吸季节变异性与降雨频率和降雨强度密切相关,在降雨量一定的情况下,较低的降雨频率和较高的降雨强度会增加土壤呼吸的变异性。呼伦贝尔草甸草原而言,在生长季土壤平均含水量为16.5%时,土壤呼吸的温度敏感性值( $Q_{10}$ )为2.12;而平均土壤含水量为26%时, $Q_{10}$ 值为2.82,明显高于前者,土壤含水量与 $Q_{10}$ 之间存在正相关关系。降雨导致土壤呼吸的激发效应和抑制效应交替发生,使草地土壤呼吸的季节变异性增加,降雨格局变化必然会对草地碳循环和碳通量特征产生深刻影响。

**关键词:**土壤呼吸;降雨;季节变异; $Q_{10}$

## Effect of rainfall on the seasonal variation of soil respiration in Hulunbe Meadow Steppe

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**Abstract:** As the main determinant of carbon balance, soil respiration is one of the major pathways for carbon loss from the soil to atmosphere and is largely influenced by soil temperature and soil moisture. During the growing season, discrete rainfall events create an important disturbance to soil respiration and lead to large uncertainties in estimating carbon exchange, especially in arid and semiarid regions. In order to examine the effect of rainfall events on soil respiration, measurements were conducted in natural rainfall conditions at a field site in the Hulunbe meadow steppe. Soil respiration was continuously monitored by an automated chamber system (LI-8150) during the growing seasons from 2009 to 2012. Five chambers (diameter 20 cm) were set in the field with at least 10 m separation between them. The results showed that the response of soil respiration to rainfall displayed different patterns, including the “birch effect” and “inhibiting effect”. Soil respiration doubled in the 1—2 hours during or after rainfall. Once rainfall reached 7—8 mm or the soil water content reached 29%—30%, which was almost equal to the value of field moisture capacity, soil respiration would be inhibited and noticeably decreased. It was suggested that moderate rainfall may stimulate the roots and microbial activities, thus enhancing soil respiration. However, heavy rainfall or an excess of moisture could reduce the soil air-filled pore space and increase anaerobism, suppressing soil CO<sub>2</sub> flux. The increase in diurnal average soil respiration was often expressed on the day following rainfall when the “birch effect” occurred, but decreased on the day of rainfall when the “inhibiting effect” was

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observed. The response of soil respiration to the “inhibiting effect” seemed quicker than the response for the “birch effect”. It was explained that the “inhibiting effect” is mainly a result of a rapid physical process and the “birch effect” was controlled by a slow biological process. Compared to the amount of rainfall, the frequency and density of rainfall were much more important factors in determining the seasonal variation of soil respiration. Over a certain amount of rainfall, a lower frequency and higher density of rainfall could strongly enhance the variation of soil respiration. During the growing seasons from May to September, the temperature sensitivity of soil respiration ( $Q_{10}$ ) in the Hulunber grassland was 2.82, when the average soil water content was 26%. This was higher than that for a soil water content of 16.5% where  $Q_{10} = 2.12$ . The  $Q_{10}$  had a positive relationship with soil water content. These results demonstrated that rainfall events had a significant influence on soil respiration and strongly enhanced its seasonal variation, which depended not only on the amount of rainfall, but also on its frequency and density. The seasonal distribution of rainfall may have important consequences for the estimation of carbon flux and C balance.

**Key Words:** soil respiration; rainfall; seasonal variation;  $Q_{10}$

土壤呼吸作用是土壤有机碳输出的主要途径,是生态系统碳循环的重要环节之一<sup>[1]</sup>。土壤呼吸作用的主要环境控制因子是温度和水分<sup>[2-3]</sup>,而降雨是土壤水分的主要来源,对地下物理、化学和生物过程具有重要的调控作用<sup>[4]</sup>,包括土壤温度、通气状况<sup>[5]</sup>、土壤微生物和根系的生理活性<sup>[6]</sup>以及呼吸底物的组成与有效性<sup>[7]</sup>等,这些关键要素直接影响土壤释放二氧化碳的过程,因而降雨发生必然对土壤呼吸作用产生重要影响。早在1958年,Birch就观测到降雨会强烈地激发土壤呼吸,这一现象被称为“Birch效应”<sup>[8]</sup>,这种效应在干旱和半干旱地区尤为明显<sup>[9]</sup>。同时,降雨量过大也会对土壤呼吸产生抑制作用<sup>[10]</sup>。这种降雨扰动引起的土壤呼吸激发效应和抑制效应,会进一步影响到草地土壤呼吸时空特征以及碳通量估算的准确性,因此,研究降雨对草地土壤呼吸变异性的影响尤为重要。本文以呼伦贝尔草甸草原为研究对象,通过对自然降雨条件下土壤呼吸的野外连续观测,研究生长季不同降雨强度和格局对土壤呼吸特征及其变异性的影响,为探讨气候变化背景下草原生态系统碳循环对降雨格局变化的响应提供科学依据。

## 1 材料和方法

### 1.1 研究区概况

研究区呼伦贝尔草甸草原地处大兴安岭西麓平缓丘陵地带,属于温带半干旱大陆性气候,气候特点是冬季寒冷漫长,夏季温凉短促,春季干燥风大,秋季气温骤降霜冻早。年均气温-5—0℃,昼夜温差和年温较差大,最冷月(1月)平均气温在-18—30℃之间,最热月(7月)平均气温在16—21℃之间,大于10℃积温1780—1820℃,无霜期85—155 d;降水量变率大,分布不均匀,年际变化也大,年降水量250—350 mm左右,且多集中在6—9月,占全年降水的75%。代表性植被类型为温带草甸草原,地带性植被主体为线叶菊草甸草原、贝加尔针茅草甸草原、羊草杂类草草甸草原,分别分布于丘陵上、中、底部,其中贝加尔针茅草甸草原为该区域典型代表类型。土壤类型主要为黑钙土或暗栗钙土。

本研究依托呼伦贝尔草原生态系统国家野外科学观测研究站进行,试验站位于内蒙古自治区呼伦贝尔市海拉尔区谢尔塔拉镇,站区地理位置N 49°19', E 120°03', 海拔约628 m。试验选择位于谢尔塔拉11队的贝加尔针茅草甸草原观测样地,面积约3.34×10<sup>4</sup> m<sup>2</sup>,于2006年底围封,围封前为自由放牧状态,主要放牧家畜为奶牛,样地主要优势种有贝加尔针茅(*Stipa Baicalensis*)、羊草(*Leymus chinensis*)、细叶白头翁(*Pulsatilla turczaninovii*)、寸草苔(*Carex duriuscula*)等。

### 1.2 研究方法

试验观测在2009—2011年进行,土壤呼吸测定采用LI-8150土壤呼吸自动测量系统(LICOR公司,美国)进行长期连续自动观测,测量频率为1次/2 h(2009年设定为1次/1 h)。该测量系统由红外分析仪主机、多路器和长期观测气室(直径20 cm)组成,本研究随机设置了5个长期固定气室,气室间距保持至少10 m,固定土壤环插入土中约3—5 cm,确保气室测量期间的封闭性,并定期剪去气室内绿色植物地上部分,保持环内土壤和凋落物的自然状态。同时,利用附近自动气象站连续观测空气温湿度、气压、太阳辐射、风向风速、6个层次(5、10、15、20、40、80 cm深度)的土壤温度、3个层次(5、10、20 cm深度)的土壤含水量以及降雨量等环境指标的测定,测量频率为半小时1次。

本试验选择生长季5—9月自然条件下发生,且之前至少连续3 d未发生任何降雨的典型降雨事件为案例,同时对土壤呼吸速率和环境因子进行同步野外连续观测,研究不同强度和频度降雨事件对土壤呼吸的影响。

### 1.3 数据处理和分析

本文利用EXCEL进行绘图和相关回归分析。

## 2 结果与分析

### 2.1 降雨对土壤呼吸日变化的影响

降雨对土壤呼吸存在明显的激发和抑制效应,以2012年5月29日22:00至翌日14:00期间发生的降雨事件为例(图1),降雨持续17:00,总降雨量达16.6 mm,平均每小时约为1 mm。降雨发生后,尽管土壤温度降低,但土壤呼吸速率逐步升高,翌日02:00达最高,由降雨前的 $1.43 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ 升高到 $2.60 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ,土壤呼吸速率增加了约1倍,此时累积降雨量达7.8 mm;之后土壤呼吸速率逐步降低至 $1.73 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ,翌日9:00后土壤呼吸速率又随土壤温度升高而增加。这说明低强度降雨对土壤呼吸存在激发作用,当降雨量累积超过一定阈值后,产生抑制作用。

而2010年7月17日18:00—20:00发生了1次短暂的高强度降雨事件,3 h内降雨量达49.8 mm,土壤呼吸速率从降雨前的 $6.34 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ 迅速降低到 $0.32 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ,降幅达95%,而表层10 cm土壤水分含量由18.5%迅速提高到29.5%。翌日14:00—16:00再次发生了6.4 mm的降雨事件,土壤呼吸速率又从 $7.40 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ 迅速减少到 $1.09 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ,降幅达85.3%,土壤含水量也由26.8%提高到29.3%。这说明短时间高强度的降雨事件对土壤呼吸表现为抑制作用,土壤水分含量较高时,较低强度的降雨事件也容易导致抑制作用发生。

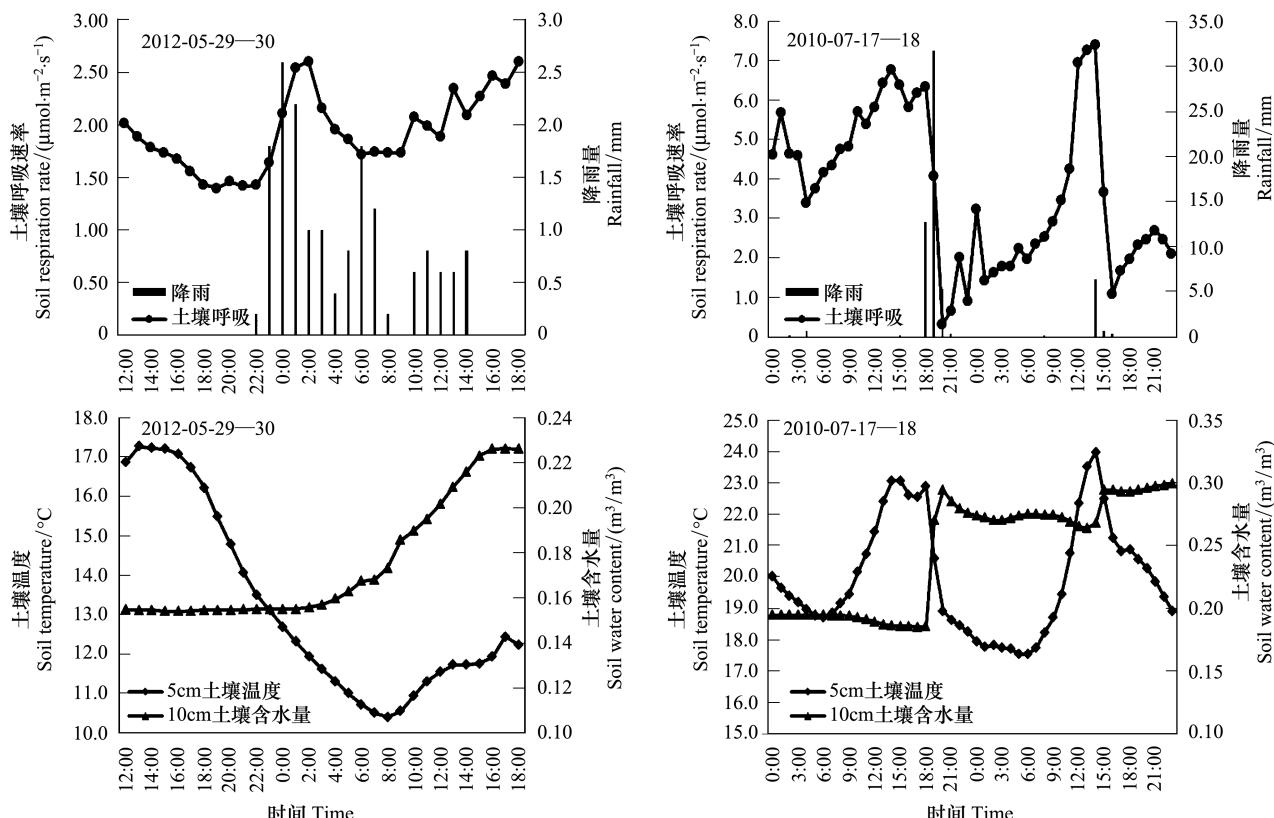


图1 降雨发生后土壤呼吸日变化

Fig.1 Diurnal variation of soil respiration with rainfall

### 2.2 降雨对土壤呼吸季节变化的影响

从图2可以看出,降雨对土壤呼吸的激发效应往往体现在次日,表现为次日平均土壤呼吸速率的显著升高;而抑制效应则在当日可以体现出来,表现为当日平均土壤呼吸速率明显降低。当首日降雨量或连续降雨累积量大于7—8 mm范围时,土壤呼吸开始呈现明显的抑制效应,且降雨量越大,抑制作用愈明显。如2009年8月20日和2010年7月17日发生的强降雨事件,单日降雨量分别为28.3 mm和50.9 mm,导致土壤呼吸速率急剧下降,降幅分别达61.5%和48.4%。

另外,降雨对土壤呼吸季节变异性也存在显著的影响。表1结果显示:2010年生长季观测期土壤呼吸速率变异系数大于2009年。2010年观测期降雨量为233.2 mm,低于2009年的249 mm,但2010年平均降雨强度(总降雨量/降雨天数)为6.1 mm/d,明显高于2009年的4.2 mm/d;降雨频率则是2010年小于2009年。这表明,土壤呼吸季节变异性大小与总降雨量关系不大,而与降雨频率和降雨强度密切相关,较低降雨频率和较高降雨强度会导致土壤呼吸的变异性增大。

### 2.3 降雨对土壤呼吸温度敏感性的影响

呼伦贝尔草原而言,降雨主要集中在生长季,全年降雨的65%—70%集中在6—9月份<sup>[11]</sup>,降雨提高了草地土壤含水量。

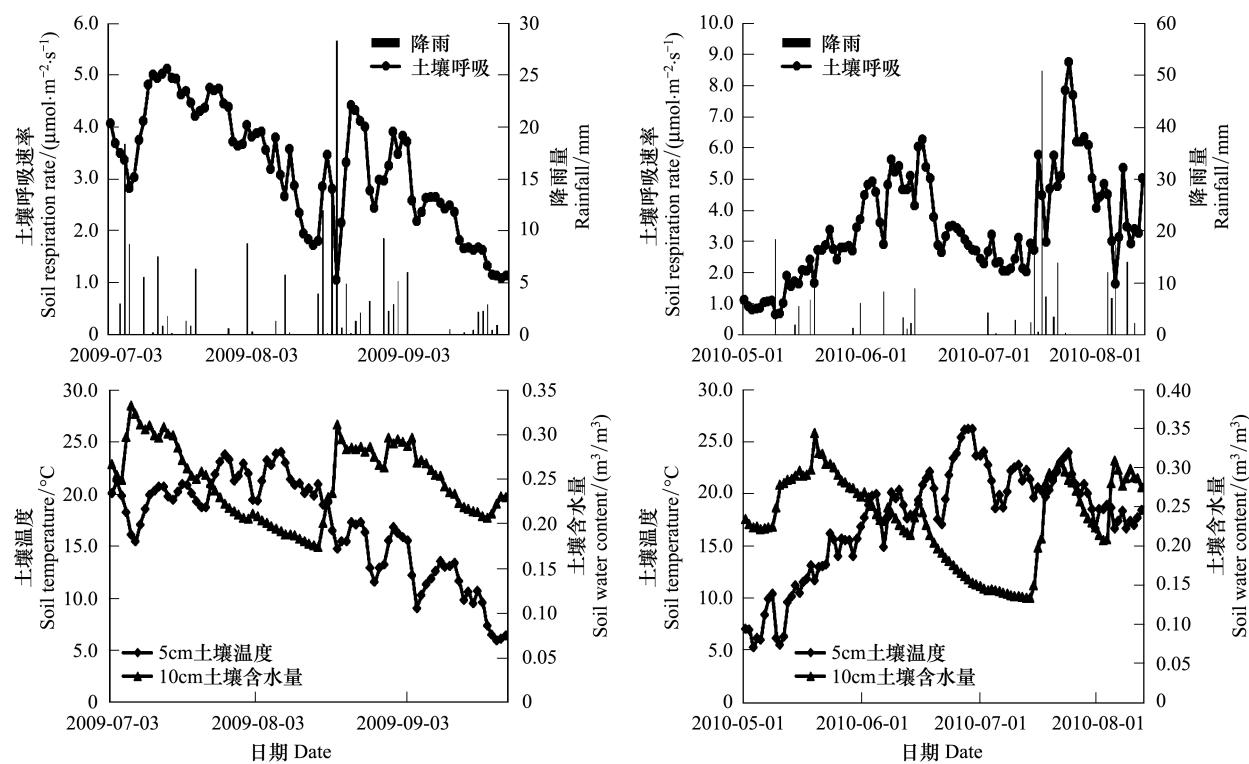


图2 降雨事件与土壤呼吸的季节动态

Fig.2 Seasonal variation of soil respiration with rainfall

表1 2009、2010年观测期间土壤呼吸速率变异系数和降雨特征

Table 1 The C.V (coefficient of variation) of soil respiration and rainfall in 2009, 2010

观测年份 Year	土壤呼吸变异系数 C.V of soil respiration	降雨量/mm Rainfall	观测天数/d Monitoring day	降雨天数/d Rainy day	降雨频率 Frequency of rain	平均降雨强度/(mm/d) Average rainfall intensity
2009	0.42	249	116	60	0.52	4.2
2010	0.49	233.2	104	38	0.37	6.1

根据近4年的气象观测数据统计,非生长季10 cm土壤平均含水量仅为11.4%,而生长季5—9月土壤平均含水量为22.2%,是非生长季土壤含水量的约2倍,降雨是生长季土壤水分的主要来源。为了分析降雨对土壤呼吸的温度敏感性影响,以平均土壤含水量22.2%为阈值,分析土壤呼吸与土壤温度的关系,由图3可以看出,土壤含水量大于22.2%时(实测点平均土壤含水量26.0%), $Q_{10}$ 值为2.85;土壤含水量低于22.2%时(实测点平均土壤含水量16.5%)的 $Q_{10}$ 值为2.12,土壤含水量越大, $Q_{10}$ 值越高,这也验证了其他研究认为土壤含水量与 $Q_{10}$ 之间存在正相关关系的结论<sup>[12-13]</sup>。

### 3 讨论与结论

降雨对土壤呼吸作用的影响包括激发效应和抑制效应,已有研究对其发生现象和机理作了很多阐释<sup>[14-19]</sup>。降雨发生后,土壤孔隙中的空气和CO<sub>2</sub>被雨水挤压外溢,短时间内(1—2 h)造成土壤释放CO<sub>2</sub>量陡增,引起“激发效应”,且降雨强度越大,这种效应越明显。对干旱土壤来说,适量的降雨改善了土壤水分状况,促进土壤团聚体的有机体底物解吸附,微生物数量和根系活性增加<sup>[8]</sup>,从而导致土壤呼吸速率增加;由于微生物和根系活性激发过程则需要一定的时间迟滞,因此,土壤呼吸的激发效应往往体现在次日,表现为次日平均土壤呼吸速率的显著升高。

本研究中,当一次或者连续降雨累积量超过一定阈值(7—8 mm)后,或土壤体积含水量大于29%—30%(与该区域土壤田

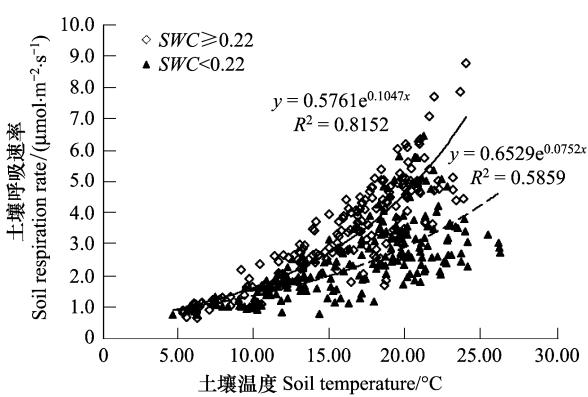


图3 不同土壤含水量(SWC)下土壤呼吸与土壤温度的关系

Fig. 3 Relationships between soil respiration and soil temperature in different soil water content

间持水量相当)时,降雨对土壤呼吸会产生明显的“抑制效应”。其因在于过量的雨水充满土壤孔隙后,土壤通透性变差,不但阻止了CO<sub>2</sub>释放途径,而且引起土壤中氧气供应减少甚至缺乏,从而抑制了土壤微生物和植物根系呼吸<sup>[20]</sup>,导致土壤CO<sub>2</sub>释放量的急剧减少,这一过程作用时间较短,甚至发生在数小时之内,直接影响了当日土壤呼吸速率的平均水平,表现为观测当日平均土壤呼吸速率的明显下降。

本研究表明,土壤呼吸季节变异性与总降雨量关系不大,而与降雨频率和平均降雨强度密切相关,在降雨总量一定的情况下,较低的降雨频率和较高的降雨强度会导致土壤呼吸变异性增大(表1),这种因降雨格局变化导致土壤呼吸的变异性增大,可能会降低仅以温度和水分为主导因子解释土壤呼吸变异性的可靠性和准确性。另外,由于土壤含水量与Q<sub>10</sub>之间存在正相关关系,降雨发生提高了土壤水分含量,使草原土壤呼吸的温度敏感性增加,这是也是导致土壤呼吸季节变异性增大的主要原因之一。因此,在利用温度-水分因子经验模型估算土壤碳排放时,需充分考虑草原生态系统降雨格局变化对土壤呼吸作用变异性的影响。

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