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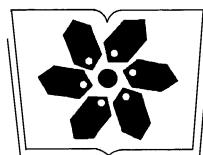
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封面图说: 永兴岛海滩植被——永兴岛是中国西沙群岛的主岛, 也是西沙群岛及南海诸岛中最大的岛屿。国务院2012年6月批准设立的地级三沙市, 管辖西沙群岛、中沙群岛、南沙群岛的岛礁及其海域, 三沙市人民政府就驻西沙永兴岛。永兴岛上自然植被密布, 野生植物有148种, 占西沙野生植物总数的89%, 主要树种有草海桐(羊角树)、麻枫桐、野枇杷、海棠树和椰树等。其中草海桐也称为羊角树, 是多年生常绿亚灌木植物, 它们总是喜欢倚在珊瑚礁岸或是与其他滨海植物聚生于海岸沙滩, 为典型的滨海植物。

彩图提供: 陈建伟教授 北京林业大学 E-mail: cites.chenjw@163.com

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植被对多年冻土的影响研究进展

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摘要:作为冰冻圈的主体,多年冻土是岩石圈与大气圈水热交换的产物,它的存在、分布及水热过程受到多种时空尺度环境因子的控制和影响。植被是生物圈的重要组成部分,是岩石圈与大气圈热量交换的媒介,它的存在和变化影响着多年冻土的水热过程和空间分布。文章综述了近几十年来植被对多年冻土影响的研究。首先,植被参与地气之间的水热周转过程,通过反射太阳辐射、贴地植被的吸水保水作用以及截留积雪作用等,对下伏冻土产生错综复杂的影响。但是不同植被类型的反射太阳辐射能力、保水与截留能力等各不相同,产生的影响大小有别。其次,同一植被类型的不同层次(如乔木层、灌木层等)对多年冻土的影响也不同,其中贴地植被产生的影响最显著。植被截留积雪,使得地面接收的太阳辐射和地表水分重分配复杂化,从而间接地影响多年冻土环境。因此当植被发生扰动后(如森林火灾和砍伐植被),就会对其所处的多年冻土环境产生各种不利影响,引发冻土灾害,甚至导致多年冻土消融。实际上植被与冻土同为寒区自然生态系统和环境的重要组成部分,它们在长期地质和生物演化中形成生态平衡。因此,植被还常常被用于指示多年冻土及其空间分布。最后提出目前研究中存在的问题,并对未来研究方向进行了展望。

关键词:植被;多年冻土;水热过程;指示物

Influences of vegetation on permafrost: a review

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Abstract: As the main body of the earth cryosphere, permafrost is formed as a result of heat-moisture exchanges between the lithosphere and atmosphere. The occurrence, distribution and thermal regime of permafrost are affected by many environmental factors of varied spatiotemporal scales. As the medium, a participant and a buffer layer of heat-moisture exchanges between the lithosphere and atmosphere, vegetation and its changes significantly affect the hydro-thermal regimes and distributive features of permafrost. In this paper, the influences of vegetation on permafrost were systematically reviewed. Vegetation has relatively high albedo and capability for water holding and retention, which vary with different types of vegetation and have great impacts on the heat-moisture regimes of permafrost. In vertical structure, there are distinct layers for a given type of vegetation. Each layer has different influences on the thermal regime of permafrost, especially the ground cover layer. Compared to the area above crown canopy, which usually reflects and absorbs probably 40% of solar radiation, the crown canopy significantly affects net radiation and ground surface temperatures in forest. Vegetation, especially the canopy structure, considerably controls the snow accumulation and redistribution through the

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interception component of the water budget, wind speed, and solar radiation at the snow surface, resulting in a significant impact on the heat-moisture regimes of permafrost. Therefore, deforestation, such as for farming or fuels, may alter snow accumulation, albedo, surface coverage, and capability of water holding of vegetation. This may greatly modify the permafrost environment, induce many types of permafrost hazards, or even result in the thawing and eventually decaying of permafrost patches. Permafrost and vegetation are key components of ecological system in cold regions. They are interdependent and interactive, and have reached dynamic equilibria in the geological and biological evolutionary processes. In permafrost regions, specific vegetation types or their combinations usually present because of appropriate temperatures and moisture conditions. Therefore, the occurrences of specific plant species or a plant community can largely indicate the occurrence of permafrost. Finally, some problems unresolved in the research at present are identified in this paper, and the prospect for further study is also proposed.

Key Words: vegetation; permafrost; heat-moisture exchange; indicator

多年冻土指的是地表下一定深度内地温持续两年以上处于0 °C以下的土层(土壤、土和岩石),它是地质历史和气候变迁背景下受区域地理环境、地质构造、岩性、水文和地被特征等因素共同影响下通过地气间物质和能量交换而发育的客观地质实体^[1],属于地球冰冻圈系统中的主要组成部分。其存在、分布以及水、热、质状态受到多种因素的影响。在全球和大陆尺度上,冻土主要受气候的纬度地带性规律的控制^[2-3];在区域尺度上,主要受海拔高度和经度(大陆度/干燥度)的控制和影响;而在局域尺度上,各种局地因子,如植被、积雪、坡度、坡向、岩性、水分等,对冻土的分布和其它特征的影响不断增强^[4-5]。其中,植被对多年冻土的水热过程及空间分布的影响非常显著,一直以来受到许多专家和学者的关注。

植被作为生物圈的重要组成部分,是大气圈和岩石圈之间热量交换的媒介,因此,植被变化可以对下伏冻土产生直接影响。例如,沼泽湿地具有隔热和蓄水功能,它能减缓活动层的能量交换过程,使下伏多年冻土受到不同程度的保护。一旦萎缩就会加快活动层的能量交换,造成地温升高、活动层加深、多年冻土退化。多年冻土退化后,土层的隔水功能减弱,地下水位下降,表层蓄水能力减弱,很难再维持沼泽湿地的原状,一般演化为草甸,大大减弱了对多年冻土的保护和增生作用^[6]。由此可见,植被发生变化会改变土层中的能量交换过程,直接影响多年冻土的水热状况;尤其是植被类型发生改变,可能会决定多年冻土的存在或者消失。

植被不仅类型繁多,而且结构复杂,特别是森林的林层结构,这样使得同一植被类型内部不同林层与下伏冻土之间的水热交换过程差异性较大,因此,不同林层对冻土水热过程所产生的影响也不相同。另外,植被冠层能够截留降雪和降低林下风速,影响林下积雪的累积和消融过程,从而对冻土的水热过程产生间接影响。可以看出,在人类活动日益频繁和剧烈的今天,如果铲除或者砍伐(包括间伐)植被,或者发生森林火灾等扰动行为都将对下伏冻土产生重要影响。因此,进一步研究和探讨植被对多年冻土的影响对保护冻土环境、资源开发和工程建设等具有重要的理论指导意义。

1 植被对多年冻土水热过程的影响

1.1 植被类型对多年冻土水热过程的影响

多年冻土区的地表植被类型多种多样,对下伏冻土水热过程的影响也是各不相同。常晓丽等发现大兴安岭多年冻土区杜香-真藓-落叶松林、真藓-落叶松林、塔头-落叶松林、柴桦落叶松林、塔头湿地5种典型植被类型的反射太阳辐射及塔头苔草的根系吸水能力不同,大大地影响了月平均地表温度和活动层冻融过程^[7]。祁连山大通河源区的多年冻土地带,当植被类型发生高山稀疏植被-高寒草甸-高寒沼泽草甸的空间变化时,冻土地温和土壤含水量都逐渐升高^[8]。南极Signy岛冻土地区簇花石萝地衣(*Usnea*)下的年平均地面温度为-1.9 °C,柳叶藓(*Sanionia*)为-2.6 °C^[9],并且苔藓和南极发草下的活动层厚度分别为57 cm和227 cm^[10]。阿拉斯加中部,研究发现白杨-白桦林和草地的土壤排水条件较好,夏季30 cm处的地温分别为10—12 °C和8—10 °C,而矮小黑云杉林则排水较差,地温为3 °C左右^[11]。一般来说,排水条件差的地方有利于苔藓的生

长,形成厌氧环境,这样能够促进土壤表层有机物的积累和泥炭层的发育。研究表明,有机物和泥炭层可以减缓夏季太阳辐射对地表的加热,冬季则由于冻结后导热系数的增大而导致地面热量大大散失。当冬季的放热大于夏季的吸热时,土层的年平均地温将低于0℃,非常有利于多年冻土的形成、保存或者加积^[12]。并且有机物层和泥炭层的存在会使地表温度与多年冻土上限处的温度之间会出现一个差值,此现象又称为“热补偿”或“热半导体”效应^[13-14]。热补偿现象的存在,使年平均地表温度为正值的地区,即通常认为不可能有多年冻土的地区,也有可能生成和保存多年冻土^[15-16]。可以看出,多年冻土地区沼泽湿地对冻土水热过程的影响是最大的,它的存在在一定程度上决定了冻土的存在及状态。因此,沼泽湿地是未来环境保护、资源开发和工程建设等人类活动首要考虑的植被类型。

1.2 林层结构对多年冻土水热过程的影响

植被冠层对太阳辐射(尤其是直射辐射)具有反射和遮挡作用,减小了到达林下地表的净辐射,阻滞了地表温度的变化,对冻土水热过程产生直接的影响^[17-18]。根据周梅在大兴安岭森林生态系统定位研究站的观测结果,6月中上旬植被冠层下部的净辐射通量仅为植被冠层上部的60%,将近40%左右的太阳辐射被植被冠层反射和吸收^[19]。北美寒带针叶林区,由于冠层反射和遮挡太阳辐射以及截留积雪,使得到达地面的热量减少,林间冷气流比无植被区更容易向下传输,林下年平均地表温度更低^[20]。但是在俄罗斯北部地区由于杂草-苔藓层和灌木的存在,土壤年平均地表温度提高1—2℃^[21]。周幼吾等认为植被冠层的这种作用还与坡向、坡度、植被类型(森林、灌丛和沼泽)有关^[22]。Cannone认为被厚度、结构、盖度和密度是决定植被隔热效应的重要因素^[23]。童伯良等也指出当植被覆盖度为80%—90%,厚度为4—5 cm时,可使唐古拉山南麓的地表温度年较差改变1.8—2.1℃,季节冻结与融化深度可增减0.1—0.3 m^[24],但是在环北极地区,无植被下的活动层反而更浅^[25]。因此,植被冠层反射遮挡太阳辐射产生的效应具有很大的不确定性,与植被类型、地理位置(经纬度)、地貌部位等都有一定的关系。

林下贴地植被(如地衣、苔藓等)和枯枝落叶层对冻土水热过程的影响也非常 important。在堪察加半岛零星多年冻土区开放的落叶松林中,如果贴地植被是泥炭藓,冻土上限仅为20—40 cm,森林青苔下的冻土上限为60 cm左右,矮灌丛-地衣下更深^[26]。堪察加半岛中部的亚寒带森林中,研究发现泥炭藓的导热率为0.07—0.19 W/(m·K),偃松落叶松林下枯枝落叶层(13—20 cm)的导热率0.06—0.09 W/(m·K),导致这里夏季地温变化不大、活动层只有几厘米^[27]。实际上苔藓、地衣等贴地植被的吸水保水能力非常强,能够对冻土的含水量和导热性能造成直接的影响^[28]。因为水的比热是4.186 kJ/(kg·℃),比矿质土大4—5倍,在其他条件完全相同时,饱水的苔藓地衣能使地面保持更低的温度和更浅的融深。最重要的是,冬季水分冻结,冰的导热系数(2.34 W/(m·K))是水的4倍,其结果造成地面放热量增大。由此可见,对于林层结构比较复杂的植被,研究其对冻土水热过程的影响是比较困难的,还需要根据实际情况和需要将其划分为不同的林层,综合考虑各层的影响才能得到更为精确的结果。

1.3 植被截留积雪对多年冻土水热过程的影响

截留降雪和降低林下风速是植被冠层重要的生态功能之一。正是因为植被冠层截留积雪和降低林下风速,直接影响了林下积雪的积累和消融过程,从而间接对冻土的水热过程产生影响^[29-30]。一般来说,植被条件越好,冠层对降雪的截留作用也越强,林下积雪厚度越小^[31]。Pomeroy等,和Hedstrom和Pomeroy研究表明,植被截留降雪量最多可达60%^[32-33];D'Eon也发现林下积雪的积累量比附近空地少40%左右^[34-35]。由于林下风速较小,减弱了林下感热和潜热通量,一系列经验性研究表明林下积雪的消融速度最低的时候比附近开阔空地要低75%左右^[36-37]。植被与林下积雪的这种敏感关系对森林地区冻土的水热动态过程影响很大^[38]。众所周知,积雪属于热的不良导体,它的存在改变了大气和地表之间的热量交换,是寒区年平均地温高于周围年平均气温的主导因素^[39],也是有些年平均气温接近于0℃或更低的地方(如阿尔泰山、西天山和小兴安岭)不存在多年冻土的主要影响因素之一^[40-41]。当积雪非常薄且反射率很高时,会导致地表温度很低;积雪厚度增加时,其隔热效应会逐渐增大,当厚度超过80 cm时,地面和大气之间几乎没有热交换^[41]。梁

林恒等在大兴安岭阿木尔地区的野外调查发现,从林区到裸地积雪厚度由21 cm增加到36 cm,与无雪区相比,林区积雪的增温效应为5.0 °C,裸地上积雪的增温效应为2.8 °C^[42]。但是积雪量大的时候,融雪季节的积雪有很强的降温和阻滞地温升高的作用。实际上,植被、积雪与冻土水热过程之间的关系是非常复杂的,探讨积雪对冻土水热过程的效应首先需要确定积雪量大小(厚度)和植被截留量大小。目前植被与积雪之间的机理作用尚不明确,植被-积雪-冻土水热过程亟待深入研究。

1.4 植被扰动对多年冻土水热过程的影响

植被扰动一般是指打破植被正常生长和自然演替过程的行为,包括人工采伐铲(剪)除、放牧和垦殖、森林火灾、森林病虫害、土地污染等等^[43]。所谓冻土环境,就是指在冻土存在空间里,影响其存在及变化所有因素的总和^[22]。植被扰动会造成物种组成、植被厚度、盖度、密度等的改变,引起一些环境因子发生变化,从而对下伏多年冻土产生影响^[44-48]。

森林火灾是一种常见的植被扰动方式。根据大量观测资料,火灾后浅层地温升高、活动层厚度增大是普遍现象^[49-52]。尽管火灾会使地温升高,活动层加厚,但随着时间的增长,这种状况会随着植被的慢慢恢复而回归到原来的状态。森林火后初始,植物尤其是草本物种迅速增多,但随着时间的推移,草本物种会在数量上、物种组成上发生很大变化,盖度也逐渐减少,灌木及乔木物种却盖度增加^[53-54]。Iawahana等比较了东西伯利亚雅库茨克过火区和天然林区的活动层状况后指出,在火灾后的起始阶段过火区由于反射率增大,净辐射量减少,但活动层厚度、地面热通量和蒸散发量却显著增加;之后随着植被的恢复净辐射量逐渐增加而地面热通量和活动层厚度逐渐减少;火灾后的5—19a,当植被基本完全恢复后,净辐射量、地面热通量和活动层厚度也基本恢复到原先状态^[55];这里有助于活动层恢复的主要是苔藓地衣等贴地层植被^[56-58]。植被焚烧后,除了使土层温度升高和融化深度加大外,往往促使多年冻土上限附近地下冰大量融化,形成热融现象^[59-61]及泥石流和碎屑坡等^[62-63]。

除了森林火灾外,另外一种重要的植被扰动方式是清除植被(如森林砍伐或垦殖)。森林砍伐将森林中较粗较高的木材伐掉,使得整片林地的遮光、保温、蒸腾作用减少,太阳辐射直接到达地表,从而对下伏多年冻土产生影响,使得地温升高、融化深度增加、冻土上限下降、土壤水分流失^[64-66]。陈亚明等研究表明大兴安岭森林采伐后,对进行小面积皆伐的林地来说,林型发生了改变,由明亮针叶松变成针阔混交林,森林覆盖率降低^[67],造成皆伐迹地0—20 cm深度的地温比天然林地高5°C左右,最大季节融化深度比天然林地增加了约10—20 cm。另外有研究表明,破坏植被对非冻土区和冻土区的影响是不一样的。库德里亚夫采夫认为在俄罗斯北方地区,清除植被会使非冻土地段的季节冻结深度大大增加,但对冻土地段的季节融化深度影响不大;在南方地区,植被对季节冻结和融化深度的影响与北方地区刚好相反^[2]。因此,评价清除植被对冻土水热过程的影响,同样需要综合考虑地理位置等要素。

此外,垦殖、放牧、城镇化及各种工程活动等也是人为扰动植被的几种方式。垦殖是人为地将森林、草原、湿地等转换为农田,从而改变了土壤的理化性质^[68-69],在原状植被消失的同时,冻土也发生退化和消失。放牧主要针对草原植被类型,一般放牧使得草地植物功能群和群落结构发生改变,土壤理化性质也发生改变^[70-71]。城镇化和工程活动等更不可避免地破坏植被,从而对冻土造成影响,使其退化和消失^[72]。

2 植被对多年冻土空间分布的影响

2.1 植被作为多年冻土的指示物

正如上述,植被对冻土水热过程有重要的影响,不过多年冻土也由于其低温、隔水等一些限制条件制约了一些植物的生长。长期以来,植被与多年冻土之间形成了一种生态平衡,这样植被常常被用作下伏冻土和融土的指示物^[73-75]。在中国大小兴安岭,落叶松-偃松林、落叶松-杜香林是多年冻土存在的典型植物群落^[76];在青藏公路南段,列氏蒿草沼泽化草甸和高山蒿草草甸是多年冻土发育的标志,而列氏蒿草、矮蒿草群丛是指示多年岛状冻土发育的主要群落,高山蒿草群丛下为季节冻土^[77]。根据H. C. 阿拉洛娃收集的有关植被和冻土相互联系的大量资料,在西西伯利亚北泰加林区,地衣型类的林型一般指示融区,苔藓型类的林型则指示

多年冻土存在^[2]。另外,植被不仅可以指示冻土发育,还可以指示地下冰的发育程度,指示季节融化深度。吴青柏等发现青藏高原上如果植被为丘状且盖度大于50%—80%时,下伏多年冻土一般为饱冰冻土或者具有含土冰层,如果植被为鳞状且盖度<50%时,一般为多冰、少冰冻土^[78]。Moskalenko 在研究了俄罗斯环极地地区植物种类频数和季节融化深度后统计得出,悬钩子、杜香、泥炭藓为建群种的泥炭地中,活动层厚度最小,最大值出现在具有裸露斑块的苔原带和地矮灌丛-地衣群落的地帶。其中 悬钩子还可作为季节融化深度的可靠指示剂(0.4—0.7 m)^[79]。

但是,植被作为指示物的作用是相对的,即在不同条件下,同一种植被可以是不同类型冻土或者是非冻土的指示物。比如,在地表无积水情况下,沼泽上以苔草和带有柳树的沼泽化草甸占优势,表明存在潜水补给。这种植被既可以指示贯穿融区(冻土层较薄),也可以指示非贯穿融区(冻土厚度很大)^[2]。因此,这种指示性是植被与冻土关系的经验统计与累积的结果,仅对一个具体地区是确定的,如果将其应用于其它地区的冻土填图,还需先深入研究该地区植物群落和冻土条件之间的联系。

2.2 植被与多年冻土空间分布

植被作为下伏冻土和融土的指示物,它的分布与多年冻土分布在空间上具有一定的相关性。戈罗德科夫研究发现在俄罗斯北方边区松林下生长的泥炭层显示不存在冻土,而丘状起伏的无林泥炭层却为冻土零星分布区所特有,表示这里接近冻土南界^[2]。在完成的 Alaska 中部 1.8 km² 的河谷区的植被和冻土分布图上,Dingman 和 Koutz 发现白云杉-桦树林可以指示冻土边界的空间分布^[80]。Nguyen 等将植被类型作为近地表面多年冻土存在的指示变量,成功预测不同植被类型覆盖下的多年冻土空间分布状况^[81]。在青藏高原冻土区,高寒植被生态系统的空间分布与冻土的分布区划同样具有很好的对应性,高寒草原分布区与高平原-河谷平原冻土区对应,高寒草甸和高寒沼泽草甸则主要集中分布于丘陵低山和唐古拉山冻土区的北麓和半阴坡^[82]。大兴安岭地区植被和冻土在空间分布上也有一定的规律性^[76]。研究表明,落叶松-偃松-泥炭藓沼泽、兴安落叶松-笃斯越桔-泥炭藓沼泽、兴安落叶松-狭叶杜香-泥炭藓沼泽等 3 种沼泽类型分区与大片连续多年冻土区、岛状或带状融区的多年冻土区、季节冻土区的实际界线具有一一对应的特征^[83]。研究发现,沼泽和多年冻土的具有共生的特征模式。当空间尺度由大、中、小变化时,共生模式依次从结构性、功能性和过程性转变;尤其是中尺度的空间分布具有高度一致性,表现在大兴安岭北部沼泽强发育区与连续多年冻土区吻合,中部沼泽中度发育区与岛状融区多年冻土区吻合,南部以及山地外围沼泽弱发育区与岛状冻土区一致,同时也说明沼泽湿地发育愈强的地区,多年冻土发育也愈厚愈稳定^[84]。总的来说,根据植被情况判定土壤冻结状态,这种指示作用对编制工程地质和多年冻土空间分布图很有帮助。如上所述,这种指示作用和多年冻土空间分布的关系多年研究工作经验的积累,对于植被-冻土关系明朗的地区,利用这种指示关系制作冻土分布图既快捷又经济。但它不具有普适性,在研究资料严重不足的地区,首先需要对植被-冻土关系进行进一步深入研究。

3 结语和展望

植被主要是通过不同植被类型和林层结果反射和散射太阳辐射、阻风挡雪、贴地植被与植物根系的吸水保水等作用来影响下伏多年冻土的水热周转过程,从而和冻土达成物质能量平衡和生态平衡。因此,这些植被一旦被破坏,有可能对多年冻土造成不可逆转的影响。另外,也正是由于植被与多年冻土之间的生态平衡和依存关系,一些植物就可以成为土壤是否冻结的指示物,指示多年冻土的类型和空间分布,不过这种指示作用具有局域性,大范围作图等的应用还需谨慎。实际上植被对多年冻土的影响是十分复杂的,植被类型(森林、灌丛和沼泽)、覆盖度、地形地貌、地理位置等因素都增加了植被对多年冻土影响结果的不确定性,因此这方面的研究还需进一步的深入和加强。关于植被对多年冻土影响的研究,目前尚存在以下问题:

(1) 多数研究比较零散,缺乏长期系统的观测系统和研究资料。据了解,国外 SNOWMIP2 (Snow Model Intercomparison Project for Forest snow Processes 2) 计划中曾在瑞士的 Alptal、加拿大的 BERMS 等的森林区对植被、积雪、冻土及其他相关内容建立了全面的观测系统^[85],并获得了长期且详实系统的数据,除此之外,其他地区则相对比较零散。因此,应该在多年冻土区建立关于植被-冻土的全面且长期有效的观测系统,进行不

同植被类型、地形地貌及不同区域之间的同步对比观测研究。

(2) 即使同一林区,植被盖度、密度和组分等在不同地方会有些许不同,能够下伏多年冻土产生差异性影响。完全依靠观测分析植被对冻土的影响,既不经济而且观测周期较长。而且如果要进行植被扰动对多年冻土影响的研究,单靠观测不仅周期长,还破坏环境。所以有必要结合合适的模型,完成各种情景的假设以及对未来情景的预测。

(3) 除植被之外,积雪也是影响多年冻土的一个重要因素。在一些积雪较长的地区,积雪对冻土的影响甚至超过植被^[40]。植被具有截留积雪的作用,但是目前关于植被、积雪、冻土之间的影响机理还不是很清楚。因此,今后无论是建立观测系统还是进行模型模拟,都应该将积雪这个因素和植被截留作用考虑进去,这样有助于得到更为精确的结果。

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